

FIRST MACROFAUNAL EVIDENCE IN THE DEVONIAN CODO FORMATION FROM ARGENTINA AND INSIGHTS INTO THE YOUNGEST CALMONIID TRILOBITE RECORDS

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ABSTRACT – We report the discovery of macrofauna in the Devonian Codo Formation, from the Precordillera of the San Juan Province, Argentina, based on a poorly preserved calmonioid trilobite. The cephalic remains resemble *Punillaspis argentina* previously reported in the late Givetian Chigua Formation, considered the lateral equivalent of the Codo Formation to the north. However, the thorax most closely resembles acastomorph calmonioids like *Pennaia* and relatives, widely distributed in the Early and Middle Devonian of Southwestern Gondwana (Malvinohosanic = Malvinokaffric Realm). Taking into account previous palynological evidence and the youngest stratigraphic records of calmonioid trilobites, the evidence supports most probably late Givetian age for the Codo Formation.

Keywords: calmonioids, trilobites, Malvinohosanic Realm, Devonian, Argentina, Brazil.

INTRODUCTION

Lower to Middle Devonian trilobites from southern South America, South Africa and Antarctica have proven to be critical for elucidating geographic distribution patterns in order to recognize a major biogeographic Devonian unit known as the Malvinohosanic (= Malvinokaffric) Realm (Clarke, 1913; Richter & Richter, 1942; Eldredge & Ormiston, 1979; Boucot & Racheboeuf, 1993; Penn-Clarke & Harper, 2021, 2023). Corresponding to cold waters of the austral circumpolar region of Southwestern Gondwana, some emblematic groups, like the family Calmoniidae Delo, 1935, are entirely endemic Malvinohosanic and display one of the most spectacular diversifications in the history of trilobites (e.g., Eldredge & Braniša, 1980; Abe & Lieberman, 2009, 2012).

However, due to the high endemism and the absence of marker groups (remarkably, graptolites and conodonts), Devonian basins of SW Gondwana lack biostratigraphic precisions based on macrofauna, and accurate stratigraphic correlations have become difficult sometimes even at regional to local scale. Palynostratigraphy has, therefore, become the main tool for supporting revised biostratigraphic schemes during the last few decades. This approach now provides a more robust framework

for studying the geographic and stratigraphic distribution of macrofauna (e.g., Amenábar & di Pasquo, 2008; Amenábar *et al.*, 2009; Grahn *et al.*, 2013; García Muro *et al.*, 2017, 2018). Thus, a number of issues have finally been resolved, such as the problem about the diachronism of the classical biozone of the brachiopod *Scaphiocoelia* in the Lower Devonian (Dalenz Farjat *et al.*, 2024).

In the particular case of Argentina, the record of trilobites has been highlighted by Baldi (1975) as relevant for developing local to regional biostratigraphic schemes, although just very preliminary insights have been considered up to date (e.g., Rustán, 2016).

The bulk of Argentine Devonian trilobites comes from the Precordillera Basin, in western Argentina, mainly from the Lower Devonian Talacasto Formation (Padula *et al.*, 1967), of the San Juan and La Rioja provinces (Baldi, 1975; Rustán *et al.*, 2011). In turn, the Middle Devonian Chigua Formation from the San Juan Province also shows relatively diverse trilobite records (Baldi, 1967; Baldi & Longobucco, 1977; Amenábar *et al.*, 2009; Allaire *et al.*, 2024).

In this context, we report a Devonian calmonioid trilobite remain from the Codo Formation (Guerstein *et al.*, 1965; Sessarego *et al.*, 1990) in the Precordillera of the San Juan Province, a lithostratigraphic unit otherwise lacking macrofaunal

records. Although very poorly preserved, this discovery allows for preliminary comparison to some trilobites of the probably equivalent Chigua Formation and the older Talacasto Formation in the Precordillera. Stratigraphic distribution of calmoniids trilobites provides new precision into ages previously considered for the Codo Formation based on palynomorphs. Some remarks on the youngest records of calmoniids trilobites, on the basis of former Brazilian evidence and the new Argentinian ones, are also provided.

GEOLOGICAL SETTING AND STRATIGRAPHY

The Argentine Precordillera corresponds to a fold-thrust belt with a preponderant eastward vergence, forming the foothills of the Andean Mountain range between 29–33°S (Ramos *et al.*, 1986; Allmendinger *et al.*, 1990). Its stratigraphy is mainly composed of a thick Paleozoic sedimentary succession, with a marked absence of volcanic and pyroclastic deposits (Furque & Cuerdo, 1979; Astini *et al.*, 1995). The recorded sedimentary fill is initially linked with the development of a Cambro–Ordovician passive margin, evidenced by the presence of a widespread carbonate platform (*e.g.*, Keller, 1999), overlain by Ordovician, Silurian and Devonian clastic sequences deposited in a foreland basin context (see Astini *et al.*, 2005; Arnol *et al.*, 2020).

The Western Precordillera (*sensu* Baldi & Chebli, 1969; Ortiz & Zambrano, 1981) is characterized by the presence of slope and platform deposits intercalated with oceanic basement volcanic rocks of Ordovician age (Boedo *et al.*, 2021). Above these, strongly deformed marine clastic successions of Silurian–Devonian age are recorded (Von Gossen, 1997; Ariza *et al.*, 2018, among others).

One of the most representative units of Devonian sedimentation within the context of the Western Precordillera is the Codo Formation (Guerstein *et al.*, 1965; Sessarego *et al.*, 1990). Outcrops of this unit occur on both margins of the San Juan River, near the town of Calingasta, where its type locality was defined along the Km 116, 117 and Del Tigre creeks (Figure 1). In the study area, the basal section of the unit does not outcrop, as the unit is in tectonic contact through the Alto de los Pajaritos Thrust with the Ordovician Don Polo Formation (Figure 1). Its top underlies in angular unconformity with Carboniferous sediments corresponding to the Del Ratón Formation (Guerstein *et al.*, 1965; Sessarego, 1988; Baldi & Peralta, 1999; Amenábar & di Pasquo, 2008; Colombo *et al.*, 2014; Gallastegui *et al.*, 2014; Rolan *et al.*, 2023a, b). The unit is composed of two main components: (i) fine to medium, and rarely coarse, sandstones of greenish to greenish-brown shades intercalated with (ii) siltstones and claystones of the same coloration (Sessarego *et al.*, 1990). One of the most significant characteristics of the Codo Formation is its high degree of penetrative deformation evidenced by folding and thrusting (Alonso *et al.*, 2008; Amenábar & di Pasquo, 2008; Gallastegui *et al.*, 2014; Rolan *et al.*, 2023a). Paleontological records found to date are scarce and limited to trace fossils (*i.e.*, *Chondrites* according to Baldi, 1970) and palynological records (Amenábar & di Pasquo, 2008). These records mainly consist of acritarchs (*Lophosphaeridium* spp. and *Cerastium* sp.)

and sparse and poorly preserved spores (*Apiculiretusispora* and *Cyclogranisporites*). The age of the unit was initially defined as Carboniferous (Guerstein *et al.*, 1965). Later, based on a fertile palynological sample, stratigraphic relationships, and the presence of Viséan age intrusions the age was referred to the Middle–Late Devonian (Sessarego, 1988; Sessarego *et al.*, 1990). Further studies on palynomorphs (Amenábar & di Pasquo, 2008) supported a Givetian to Frasnian age.

MATERIAL AND METHODS

The Codo Formation is virtually devoid of macrofossils, our exceptionally rare trilobite finding was made only after several collecting trips to the locality. The specimen was found in the middle section of the unit, exposed on the southern side of Provincial Route 149, prior to reaching the Km. 117 Creek (Figure 1), within a fine-grained muddy dark greenish-gray siliciclastic interval (GPS 31°15'14.4" S/69°19'48.12" W).

The trilobite remain is about 1.5 cm long (exsagittal) and 1 cm wide (transversely), including the left side of the cephalon poorly preserved and a few anterior (up to 7–8), outstretched, articulated, partial thoracic segments preserving light-colored exfoliated carapace (Figure 2A–F).

This specimen is housed in the collection of the Geological Research Institute Dr. Emiliano Pedro Aparicio at Universidad Nacional de San Juan, numbered as INGeo-PI-2305.

Two additional *Punillaspis argentina* (Baldi, 1967) calmonioid trilobite specimens (thoraces), coming from nodules of the upper part of the Chavela Member of the Chigua Formation at its type-area near Chigua, to the north of Malimán town, are illustrated for comparison herein (Figure 2L–N, P–R) and housed in the same collection. INGeo-PI-2306 comes from the Chavela Creek at GPS 29°53'37.5" S/69°06'14.28" W. INGeo-PI-2307 comes from the Don Justo Creek at GPS 29°54'25.8" S/69°06'25.3" W. These *Punillaspis* specimens are late Givetian in age, since they co-occur with the goniatite *Epitornoceras baldi* (Leanza, 1968), as stated by Allaire *et al.* (2024).

A couple of specimens of the acastomorph calmonioid *Pennaia verneuili* (d'Orbigny, 1842), coming from the Middle Devonian of Bolivia are also illustrated for facilitating comparisons (Figure 2G–K, O–S). They are housed in the collection of the Centro de Investigaciones Paleobiológicas (CIPAL) at the CICTERRA building, under the prefix CEGH-UNC. Both represent infaunal molts preserved in nodules, as previously reported by Rustán *et al.* (2015). The specimen CEGH-UNC 27687 probably comes from the upper Belén Formation (most probably Middle Devonian) near Belén. It is important to note that after the recognition of goniatites in the base of the Huamampampa Formation at Campo Redondo, Bolivia (Troth *et al.*, 2011), the laterally equivalent upper part of the Belén Formation yielding *Pennaia* should be considered not older than Eifelian.

CEGH-UNC 27688 comes from the overlying Sicasica Formation (Middle Devonian), near La Paz.

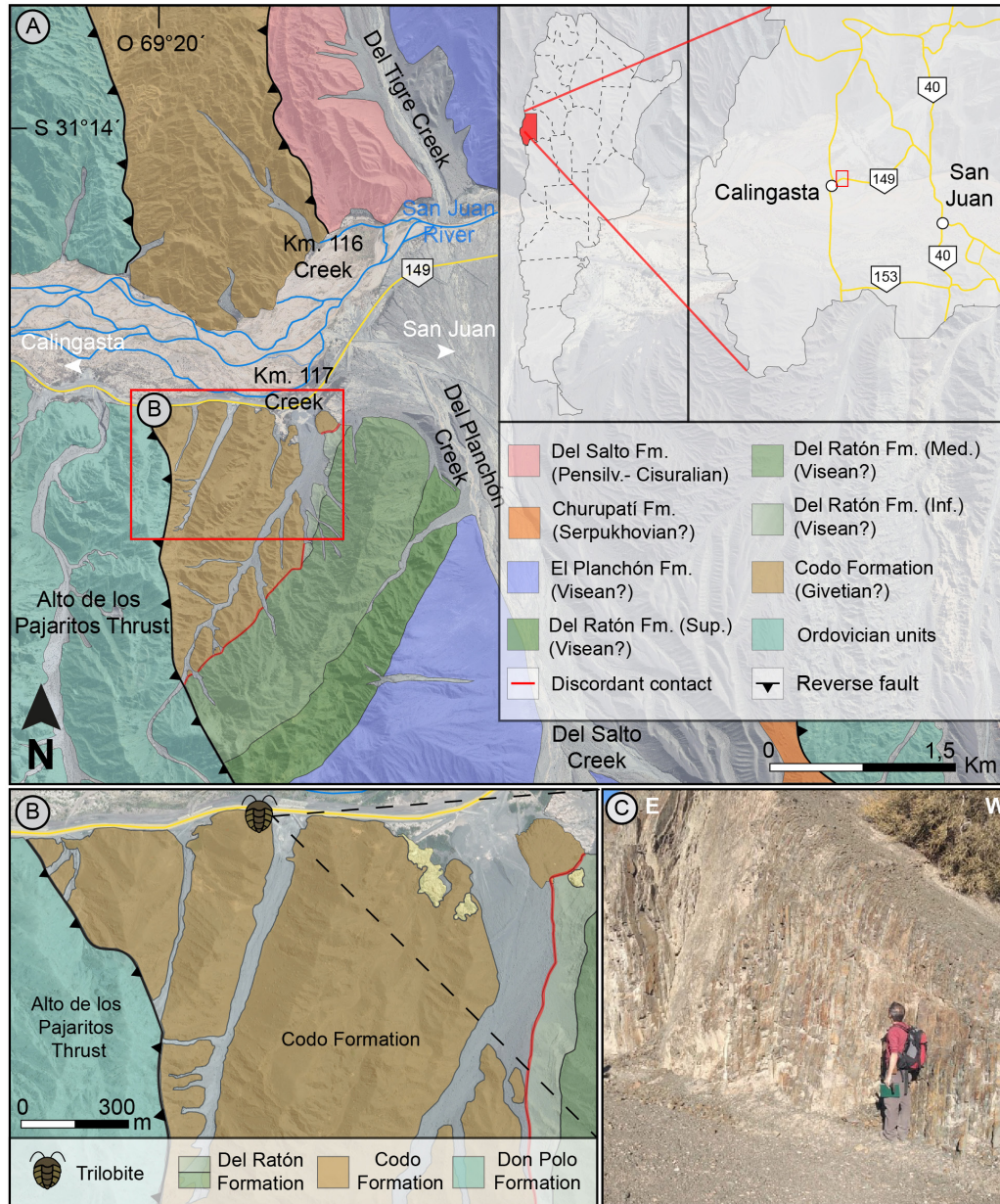


Figure 1. A, geological map and general stratigraphy of the study area located on the south margin of the San Juan River, next to the Route 149 linking San Juan and Calingasta, Argentina. B, detailed geological map showing fossil locality in the Codo Formation. C, general aspect of the Codo Formation of the fossil-bearing interval.

Fossils were prepared using pneumatic jacks and needles and then inked and whitened with ammonium chloride (NH₄Cl) sublimate, or reproduced in black latex casts of external molds, when necessary, for taking pictures under raking light using a *Canon Power Shot S50* digital camera mounted on a *Leica MZ75* binocular loupe.

SYSTEMATIC PALEONTOLOGY

Technical terms are abbreviated as traditional in trilobite descriptions: transversal/transversely (**tr.**), sagittal/sagittally

(**sag.**), exsagittal/exsagittally (**exsag.**). Morphological terminology follows Whittington & Kelly (1997).

Superfamily ACASTOIDEA Delo, 1935

Family CALMONIIDAE Delo, 1935

Calmoniidae indet.

(Figure 2A–F)

Referred material. INGEO-PI-2305, a fragmentary thorax, Codo Formation, Givetian? (Middle Devonian), San Juan Province, Argentina.

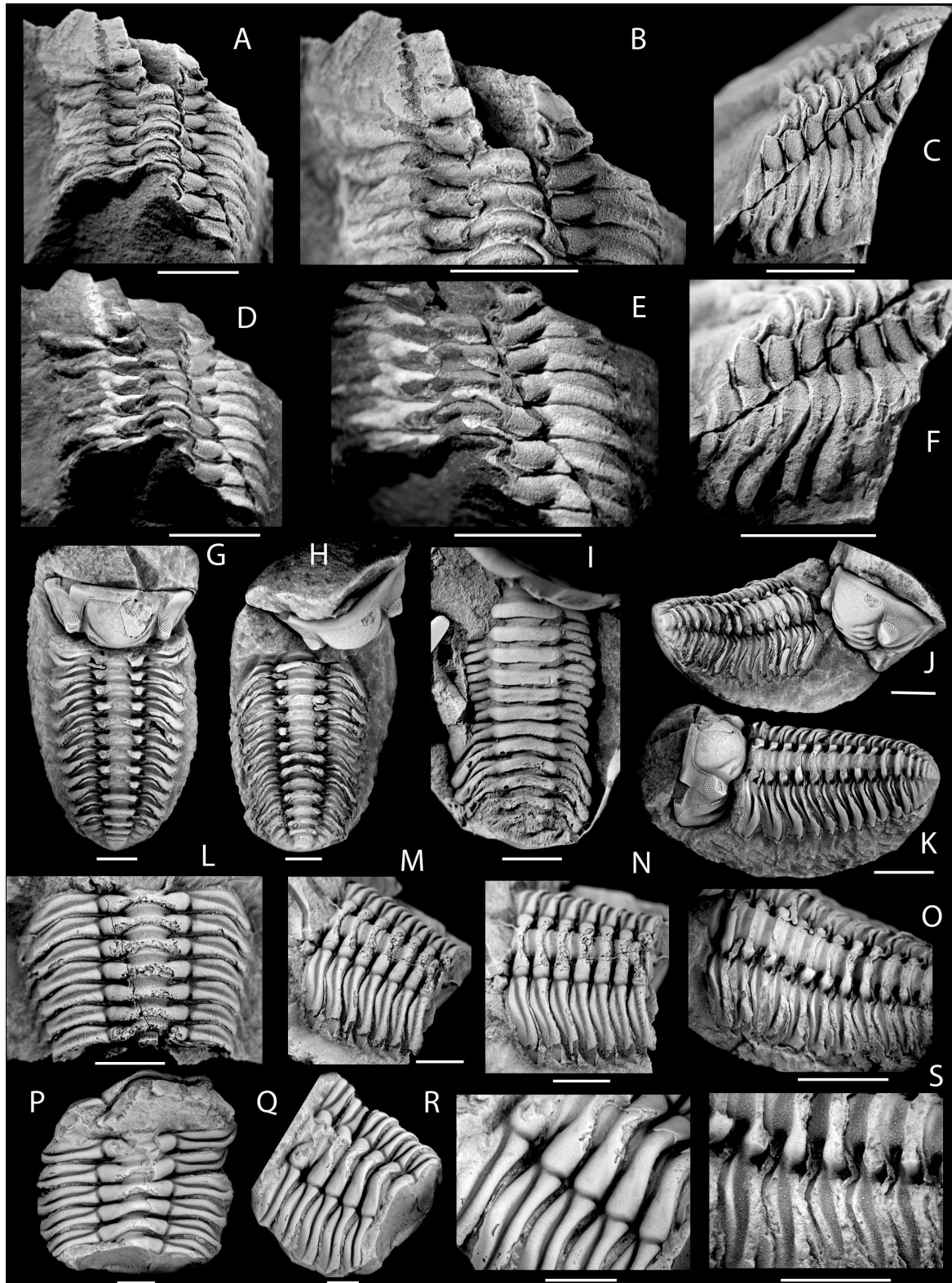


Figure 2. A calmoniid trilobite specimen, the first macrofaunal record of the Devonian Codo Formation, San Juan Province, Argentina, and some similar taxa for comparison. **A–F**, partial cephalon and thorax of the calmoniid specimen from the Codo Formation, INGEO-PI-2305; **A–E**, dorsal view (**D** and **E** without whitening in order to show exfoliation); **C** and **F** latero-dorsal view. **G–K**, **O**, **S**, the acastomorph calmoniid *Pennaia verneuli* (d’Orbigny, 1842), from the Middle Devonian of Bolivia; **G**, **K**, CEGH-UNC 27688, from the Sicasica Formation, near La Paz; **H–J**, **O**, **S**, CEGH-UNC 27687, from the upper part of the underlying Belén Formation, near Belén. **L–N**, **P–R**, the calmoniid *Punillaspis argentina* (Baldis, 1967), from the upper Givetian (Middle Devonian) of the upper part of the Chigua Formation (Chavela Member) in its type-area, San Juan Province, Argentina; **L–N**, INGEO-PI-2306, part of a thorax partially decorticated, from the Chavela Creek, **L**, dorsal view, **M–N**, latero-dorsal view; **P–R**, INGEO-PI-2307 part of a thorax preserving carapace, from the Don Justo Creek, **P**, dorsal view; **Q–R** latero-dorsal view. Scale bars = 5 mm.

DISCUSSION

Taxonomic remarks

Thoracic segments of our specimen do not suggest a trilobite referable to Phacopidae, especially when taking into account their proximal pleural region (in between the axial furrow and fulcrum). For example, each thoracic segment does not show typical phacopid features as the anterior and posterior margin of each proximal part of pleurae nearly straight (when exfoliated), with a ridge-like anterior pleural band narrower than the posterior one, and pleural furrow deeply incised. In contrast, the proximal part of each pleurae shows the anterior margin gently convex forward, anterior and posterior pleural bands of similar height and width (exsag.), separated by a shallow and wide (exsag.) pleural furrow. In addition, although poorly preserved, the distal (external to fulcrum) region of pleurae appears to be narrow and elongated (exsag.), thus differing from the wide and rounded common morphology of Phacopidae, which develops a wide paddle-like articulating facet.

Conversely, such characters, in combination with the very wide axis, the very narrow (tr.) proximal pleural region, and the distal morphology and orientation of pleurae, indicate a calmoni trilobite. Furthermore, the posterior left part of the preserved cephalon appears to show the typical S1 and L1 morphology of an acastid or acastomorph calmoni.

In addition, within calmoniids, the seemingly very wide lateral and post-ocular fixigenal area suggests minute eyes. This character is typical of the calmoni *Punillaspis argentina* (Baldi, 1967) only known from the Chigua Formation, in the San Juan Province, a lithostratigraphic unit interpreted as a lateral stratigraphic equivalent of the Codo Formation (Amenábar & di Pasquo, 2008). However, we consider that a reliable taxonomic assignment is not possible since a preservational artifact should not be excluded. Furthermore, the thorax of *Punillaspis* appears to differ from our specimen in having a proportionally narrower (tr.) axis, wider (tr.) proximal pleural region, lateral lobe of the axial ring more inflated and without an apodemal notch posteriorly located, and proximal portion of posterior pleural band adjacent to axial furrow more inflated and higher than the anterior pleural band. *Punillaspis* has, otherwise, been interpreted as a member of the derived *Metacryphaeus* group (Lieberman, 1993; Carbonaro *et al.*, 2018). For comparative purposes, we provide illustrations of some *Punillaspis* thoraxes coming from its type-area (Figure 2L–N, P–R).

Regarding other possible affinities for the new material we present, Edgecombe (1993) highlighted thoracic similarities between the basal acastomorph calmoni *Andinacaste* Eldredge & Braniša, 1980, from the Silurian–Devonian transition of Bolivia and Argentina, and the secondarily convergent acastomorph *Pennaia* Clarke, 1913, particularly in the absence of spinous processes and presence of notches in the distal tips of pleurae behind of a slender articulating facet in the last thoracic segments.

The thoracic morphology of *Andinacaste* is however not well-illustrated, particularly in the case of the Argentinian *Andinacaste espejensis* Edgecombe, Waisfeld & Vaccari, 1994, from the Ludlow of the Los Espejos Formation, San Juan Province. The

better-known thoracic morphology of *Andinacaste legrandi* Eldredge & Braniša, 1980, from the Lochkovian uppermost part of the Catavi Formation from central Bolivia, appears to differ from *Pennaia* and from our specimen in its slightly wider (tr.) axial lobe, without conspicuous lateral expansions, and pleurae with anterior pleural bands slightly shorter (tr.) and more fusiform in dorsal view. Coaptative structures, particularly nodes, and notches related to the articulating facet in the terminal portion of pleurae, cannot be observed in our specimen due to decortication so that a comparison is not warranted.

Pennaia represents another possibility. *Pennaia* (particularly species from the Devonian of Bolivia) had previously been assigned to *Acastoides* Delo, 1935. However, Eldredge & Braniša (1980) convincingly demonstrated that *Acastoides* is absent from the Devonian of Gondwana, since taxa referred to it are convergent with Silurian–Devonian Laurussian “true” *Acastoides* (also see Edgecombe, 1994, p. 56).

The taxonomic concept of *Pennaia* is somewhat vague due to the original diagnosis of Clarke (1913, p. 132). In addition, taxonomic assignments were usually supported on comparisons with similar genera, especially *Oosthuizenella* Cooper, 1982, *Phacopina* Clarke, 1913, *Calmonia* Clarke, 1913, and *Kozłowskiaspis* Braniša & Vaněk, 1973 (see Eldredge & Braniša, 1980; Cooper, 1982; Edgecombe, 1993). In some instances, assignments referred to earlier illustrations rather than to the presence of autapomorphies or diagnostic combinations of characters. The only mention of *Pennaia* in Argentina is by Waisfeld *et al.* (1994, p. 32), from uncertain levels of the Lower Devonian Talacasto Formation at Cerro El Peñón locality, San Juan Province.

All in all, the thorax of our specimen most closely resembles *Pennaia pupillus*, from the Devonian of the Bokkeveld Group, South Africa, as reported by Cooper (1982, fig. 53); *Pennaia pauliana* Clarke, 1913, from the Devonian of Brazil, and especially *Pennaia verneuili* (d’Orbigny, 1842), which is commonly found in the Pragian to Middle Devonian of the Icla, Belén and Sicasica formations from Bolivia. For a more direct comparison to *P. verneuili*, consult Edgecombe (1994, pl. 9) and see herein external and internal molds of specimens from the Belén Formation (Figure 2G–K, O–S).

The monotypic *Oosthuizenella*, from the Devonian of South Africa and Malvinas Islands, has been considered sister taxon to *Pennaia*. *Oosthuizenella ocellus* (Lake, 1904) was compared to *Pennaia* by Cooper (1982) and Edgecombe (1994) but differs from our specimen in having a more inflated lateral part of the axial ring; anterior pleural band adjacent to the axial furrow narrower (tr.) than the posterior one, more fusiform in shape; and pleural furrow narrower (exsag.), deeper and straighter.

Although *Phacopina* Clarke, 1913, is also a typical acastomorph calmoni very close to *Pennaia*, its thoracic morphology (consider *P. convexa* Eldredge & Braniša, 1980, fig. 22) subtly differs from *Pennaia* and from our specimen in having bigger apodemal notches posterior to the lateral part of the semiarticulating half-ring (visible in internal molds) and a

wider (sag.) axial ring, which also appears to be slightly more gently expanded laterally.

Calmonia is widely distributed in the Lower to Middle Devonian of South America. However, the type species *C. signifer* Clarke, 1913 from Brazil, differs from our specimen in having a proportionally wider (tr.) pleural region (*i.e.*, proportionally wider axial lobe), axial rings seemingly more gently and gradually expanding laterally, and distal pleural region not so inclined laterally external to the fulcrum.

Kozłowskiaspis Braniša & Vaněk, 1973 (particularly *K. superna*, see Eldredge & Braniša, 1980, fig. 27) differs in having a more inflated lateral part of axial rings, anterior pleural band narrower (tr.) than the posterior one adjacent to the axial furrow, and pleural furrow narrower (exsag.), deeper and straighter, more posterolaterally directed and distal pleural region not so inclined external to the fulcrum.

Hence, we preliminary hypothesize our specimen might possibly correspond to *Punillaspis* according to cephalic characters, although thoracic ones seem to resemble acastomorph calmoniids close to *Pennaia*.

Some insights into the stratigraphic range of calmoniid trilobites

Calmoniid trilobites have a confirmed Silurian to upper Middle Devonian stratigraphic range, with a single mention of a lower Upper Devonian record in Brazil. Earliest records derive from the Llandovery to Wenlock of the Lipeón Formation, northwestern Argentina (Baldis *et al.*, 1976; Rubinstein & Toro, 2006); the Ludlow of the Los Espejos Formation, western Argentina (Edgecombe *et al.*, 1994); and the Prídolí to probably Lochkovian of the El Carmen Formation from eastern Bolivia (Edgecombe & Fortey, 2000).

In turn, based on palynomorphs and stratigraphic correlation provided by Grahn *et al.* (2010, 2013), the youngest records of calmoniid trilobites derive from the Middle–Late Devonian transition from the Paraná Basin in Brazil (Carbonaro *et al.*, 2016, 2018). These records include: (i) *Metacryphaeus australis* (Clarke, 1913) from the lithostratigraphic Unit 4 of the Chapada Group (Alto Garças Sub-Basin, in the north) and also from the equivalent upper part of the São Domingos Formation (Apucarana Sub-Basin, in the south); and (ii) *Metacryphaeus tuberculatus* (Kozłowski, 1923) from the uppermost levels of the Unit 4 of the Chapada Group.

The records of *Metacryphaeus australis*, however, appear to be no younger than Givetian. On one hand, the stratigraphic position of the record of *M. australis* from the lithostratigraphic Unit 4 of the Chapada Group is lower than the interval that possibly ranges into the Frasnian according to Carbonaro *et al.* (2016, p. 561). On the other hand, following Grahn *et al.* (2013, p. 36) “The highest outcropping strata of the São Domingos Formation in the Apucarana Sub-basin are early Givetian (Sequence E) ...”. Hence, a Late Devonian (early Frasnian) age would be possible to assume only for the record of the two *M. tuberculatus* specimens derived from the uppermost levels of the Unit 4 of the Chapada Group, at the locality of Doverlândia, as reported

by Carbonaro *et al.* (2016). However, a Frasnian age has not been positively confirmed for *M. tuberculatus* in such locality but inferred from the Givetian–early Frasnian stratigraphic range of the palynomorph of the fossil-bearing interval.

One putative late Fammenian *Metacryphaeus* record has been reported from the Longá Formation, Parnaíba Basin, northeastern Brazil (see Carvalho & Ponciano, 2015). Specimens consist of a cephalon and a pygidium supposedly from the lower part of the unit cropping out at the Barreiras Farm, Elesbão Veloso street, Valença do Piauí City, Piauí State. This is the only one faunal fossiliferous locality of the unit. This record has been interpreted as dubious due to the Strunian age interpreted for the bearing levels. Therefore, we consider this locality merits further investigation in order to discard a stratigraphic misinterpretation. Consequently, we suggest that Late Devonian records of calmoniid trilobites should be considered putative or questionable for the moment.

In this biostratigraphic scenario, *Punillaspis argentina* from the Chigua Formation, western Argentina (Baldis, 1967; Baldis & Longobucco, 1977), seems to be one of the best dated youngest calmoniid occurrences. The Chigua Formation was interpreted as a lateral stratigraphic equivalent of the Codo Formation (Amenábar & di Pasquo, 2008), cropping out nearly 155 km to the north of the locality studied in this paper, to the north of Malimán town in the San Juan Province (Baldis, 1967). The *Punillaspis*-bearing levels of the Chigua Formation are located in the uppermost part of the unit (the upper part of the Chavela Member) which reliably corresponds to the upper Givetian based on the recent revision of the ammonoid *Epitornoceras baldisi* (Leanza, 1968) by Allaire *et al.* (2024).

Therefore, although Amenábar & di Pasquo (2008) suggested Givetian to Frasnian ages for the Codo Formation on the basis of palynomorphs, the new calmoniid record would suggest ages older than Late Devonian, at least for the trilobite-bearing level. We hypothesize that a (probably late) Givetian age is the most probable one for the trilobite record reported herein, supporting a correlation with the upper part of Chigua and São Domingos formations and the upper part of Unit 4 of the Chapada Group.

DATA AVAILABILITY STATEMENT

Data will be available upon request.

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AUTHOR CONTRIBUTIONS

Juan José Rustán: writing – original draft, conceptualizations, investigation, taxonomic and biostratigraphic analysis. Julián David Rolan: writing – review, editing; conceptualization, formatting. Jonatan Ariel Arnol: review, conceptualization. Norberto Javier Uriz: review, conceptualization. All authors gave final approval for publication and agreed to be held accountable for the work performed therein.

DECLARATION OF AI USE

We have not used AI-assisted technologies to create, review, or any part of this article.

ETHICS

This work fulfills collecting authorizations.

CONFLICT OF INTEREST

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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