



A POSSIBLE LUNGFISH BURROW IN THE UPPER CRETACEOUS ADAMANTINA FORMATION (BAURU BASIN, BRAZIL) AND ITS PALEOECOLOGICAL AND PALEOENVIRONMENTAL SIGNIFICANCE

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ABSTRACT – Vertebrates produce a variety of trace fossils, mostly tracks and trackways, coprolites, and burrows resulting from fossorial and subterranean habits. Burrows, particularly, tend to represent temporary or permanent shelter. Vertebrate burrows are relatively understudied in the Brazilian Mesozoic units, as well as in Cretaceous rocks worldwide. This study aims to analyze a paleoburrow in the Upper Cretaceous Adamantina Formation and to discuss its paleobiological and paleoenvironmental implications. A descriptive study was carried out on the paleoburrow morphology, following aspects such as orientation, architecture, and dimensions, and facies association analysis provided the interpretation of the paleoenvironment in which the tracemaker lived. The origin of the paleoburrow was highlighted, as well as the size of its possible producer, besides inferring the exposure time of the paleoburrow from taphonomic interpretations, which contains allochthonous bone remains. The results suggest long periods of dry conditions, as previously suggested for the much-debated Adamantina Formation, interrupted by flooding events as evidenced by facies analysis and suggested by the lungfish burrow. The possibility of a lungfish record also expands the presence of these animals to Adamantina Formation, even though more data are needed to understand the paleoecology of these formation.

Keywords: lungfish, paleoburrow, floodplain, Adamantina Formation.

RESUMO – Os vertebrados produzem uma variedade de traços fósseis vestigiais, na sua maioria pistas e caminhos de rolamento, coprólitos e tocas resultantes de hábitos fossoriais e subterrâneos. As paleotocas, particularmente, tendem a representar abrigo temporário ou permanente. As paleotocas de vertebrados são relativamente pouco estudadas nas unidades mesozoicas brasileiras, bem como nas rochas cretácicas de todo o mundo. Este estudo visa analisar o registro de uma paleotoca no Cretáceo Superior da Formação Adamantina e discutir as suas

implicações paleobiológicas e paleoambientais. Foi realizado um estudo descritivo sobre a morfologia da paleotoca, seguindo aspectos como orientação, arquitetura, e dimensões, e a análise de associação de fácies forneceu a interpretação do paleoambiente em que o produtor viveu. Foi destacada a origem da paleotoca, bem como o tamanho do seu possível produtor, além de inferir o tempo de exposição da paleotoca a partir de interpretações tafonômicas, que contém restos ósseos alóctones. Os resultados sugerem longos períodos de condições secas, como sugerido anteriormente para a tão debatida Formação Adamantina, interrompida por eventos de inundação como evidenciado pela análise de fácies e sugerido pela toca do peixe pulmonado. A possibilidade de um registro de peixe pulmonado também expande a presença destes animais à Formação de Adamantina, embora sejam necessários mais dados para compreender a paleoecologia destas formações.

Palavras-chave: peixe pulmonado, paleotoca, planície de inundação, Formação Adamantina.

INTRODUCTION

The interactions between macroorganisms and sedimentary substrates produce different biogenic structures that are good indicators of paleoenvironmental conditions (*e.g.*, Pemberton & Frey, 1982; Bromley, 1990; Genise, 2004; Bromley & Heinberg, 2006; Melchor *et al.*, 2010; Buatois & Mángano, 2011, 2018; Sedorko *et al.*, 2020). Vertebrate trace fossils usually result from locomotion (tracks and trackways), excretion (coprolites), and fossorial habits (burrows) (Voorhies, 1975), providing relevant paleoecological and paleoenvironmental information (*e.g.*, Reichman & Smith, 1990; Groenewald *et al.*, 2001; Hasiotis *et al.*, 2004, 2007; Hembree, 2010; Dentzien-Dias & Figueiredo, 2015; Mukherjee *et al.*, 2017; Martinelli *et al.*, 2019).

Paleoburrows are mostly associated with paleosol horizons, cross-cutting stratigraphic levels, and evidence specific zoological groups. These structures may provide a relative age for an ichnoassemblage due to their inferred producer and its stratigraphic relationships (Ponomarenko & Ponomarenko, 2019). Mammals are among the main present-day fossorial vertebrates, constructing complex burrows with several tunnels, branches, entrances, and chambers (Kinlaw, 1999). However, the fossil record is plenty of other examples of fossorial lineages: Paleozoic fishes and tetrapods (Hembree *et al.*, 2004; Francischini *et al.*, 2018); Permian–Triassic reptiles and synapsids (Modesto & Botha-Brink, 2010); Jurassic–Cretaceous ornithomimid dinosaurs (Martin, 2009; Krumenacker *et al.*, 2019), synapsids (Dentzien-Dias *et al.*, 2008; Riese *et al.*, 2011), and possibly crocodylomorphs (Martinelli *et al.*, 2019).

Vertebrate burrows often represent temporary or permanent shelter for survival, reproduction and protection of cubs, and social organization (*e.g.*, Kinlaw, 1999; Varricchio *et al.*, 2007; Voigt *et al.*, 2011). These subterranean systems can reveal how the structure was dug, depending on the substrate in question, and when abandoned, they might become an empty space in the environment to be occupied or even be filled by one or more events (Kent & Snell, 1994; Kinlaw, 1999; Voigt *et al.*, 2011; Kinlaw & Grasmueck 2012; Cardonatto & Melchor, 2021).

Vertebrate burrows are relatively understudied in the Brazilian Mesozoic units, and in Cretaceous rocks worldwide. In Brazil, they occur in the Permian–Triassic

Buena Vista Formation (formerly the “Piramboia Formation”, Francischini *et al.*, 2018), the Norian Caturrita Formation (Silva & Dominato, 2010), the Upper Jurassic Guará Formation (Dentzien-Dias *et al.*, 2008), the Lower Cretaceous Botucatu Formation (Manes *et al.*, 2021), and in the Upper Cretaceous Marília and Adamantina formations (Martinelli *et al.*, 2019; Silva *et al.*, 2022). In other hand, Cretaceous burrows attributed to lungfishes are common worldwide (Surlyk *et al.*, 2008; Marshall & Rogers, 2012; Fanti *et al.*, 2016), associated to its aestivation behavior that generated chambers with variable morphologies. Lungfishes’ fossils were previously reported for Bauru Group (Alves *et al.*, 2013, 2021). In this sense, this study aims to describe a new record of a paleoburrow in the Late Cretaceous deposits of the Adamantina Formation and to discuss its paleobiological and paleoenvironmental implications.

GEOLOGICAL SETTING

The intracratonic Bauru Basin (Campanian–Maastrichtian) covers an area of approximately 330,000 km² and was originated by flexural process related to the Alto Paranaíba Uplift event, after the fragmentation of the Gondwana (Fernandes, 1998; Batezelli, 2017). The basin filling can reach 300 m and records continental paleoenvironments that characterize the Bauru Group. The Araçatuba Formation represents the basal unit of the Bauru Group. It is composed of mudstones and very fine-grained sandstones, superimposed by the fine- to medium-grained sandstones from Adamantina and Uberaba formations. These strata are covered by fine- to coarse-grained, locally conglomeratic sandstones of the Marília Formation (Batezelli, 2017). The Adamantina, Uberaba and Marília formations were interpreted as distributive fluvial systems with a northeast source area and a basinward progradational pattern (Fernandes, 1998; Batezelli & Ladeira, 2016; Batezelli *et al.*, 2019). This study focuses on the Adamantina Formation (Turonian–Santonian) in the Ituiutaba municipality (Minas Gerais State, Figure 1), from where the paleoburrow described herein comes. A diversity of sauropod and few theropods dinosaurs, squamata, testudines, anurans, fish, and crocodylomorpha somatofossils have already been reported in the sedimentary beds of the Adamantine Formation. Crocodylomorpha is the highest

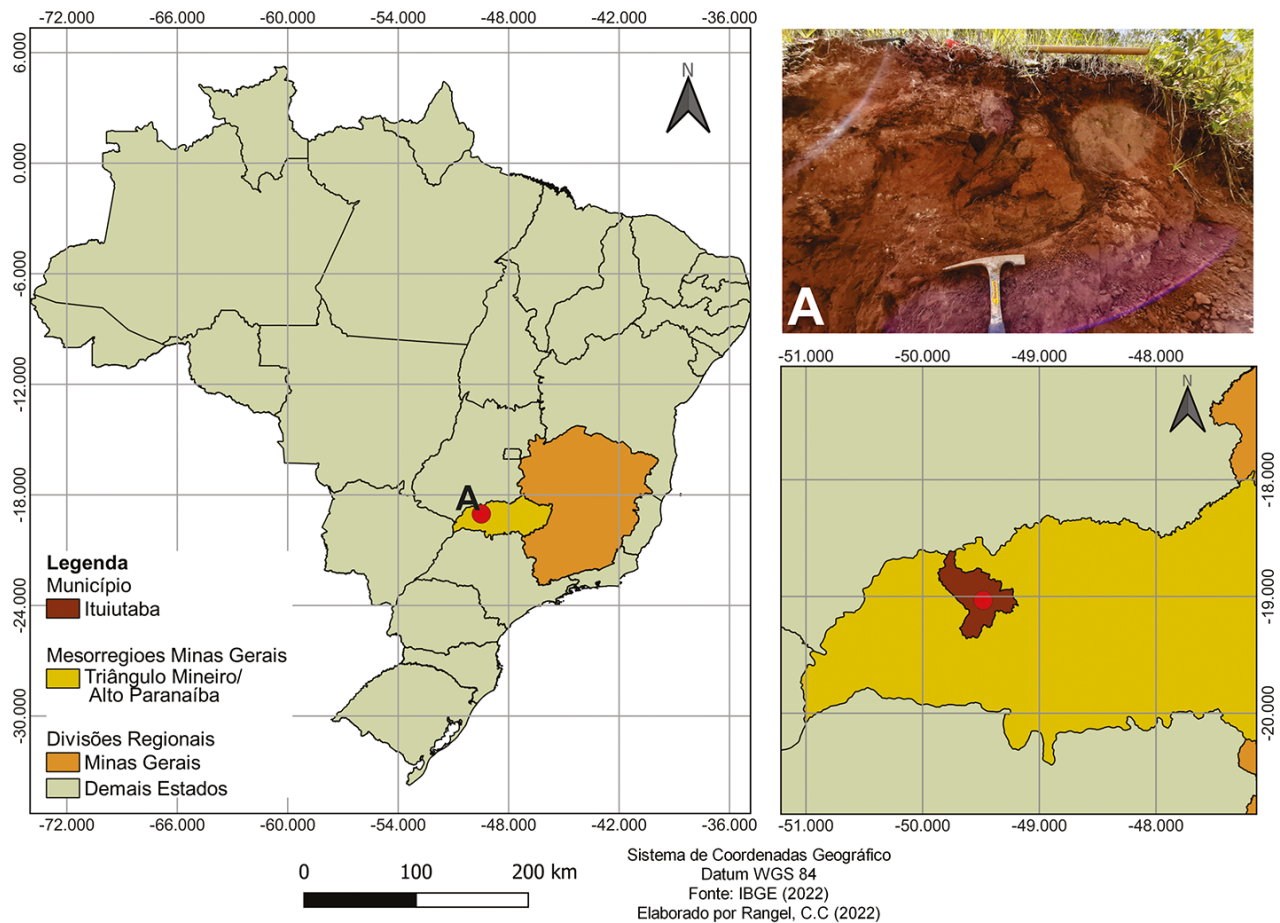


Figure 1. The paleoburrow from Brazilian Mesozoic. The green star represents the found Cretaceous paleoburrow. A, the Serra do Corpo Seco Hill from the Municipality of Ituiutaba, Triângulo Mineiro region, Minas Gerais State.

diverse group in this unit, with about 20 taxa (see Matinelli & Teixeira, 2015, Darlim *et al.*, 2021).

MATERIAL AND METHODS

The paleoburrow was found in the Serra do Corpo Seco (outcrop coordinates: 19°01'55.1"S 49°28'51.1"W), Ituiutaba Municipality, Minas Gerais State (SE Brazil). The lower part of this structure is oriented towards on the east side and the upper is perpendicular to the depositional plane. The three-dimensional paleoburrow morphology was accessed by successive excavations. Macroscopic observation, measurements with measuring tape and scale, photographs, and sketches were made during the fieldwork to carry out the paleoburrow description. The sediment fill was collected for textural characterization. The descriptive criteria followed Bordy *et al.* (2017) and Cardonatto & Melchor (2021).

The burrow was destroyed during this study, but its filling was collected. The measurements were taken on both burrow

axis, the longest and shortest. Bone fragments were found inside the paleoburrow filling and are housed at the scientific collection of the Laboratório de Paleontologia Estratigráfica-LAPE/UFU, under the number V-0126 (Figure 2). The analysis of the taphonomic aspects of these fragments considered its weathering stage, following Behrensmeier (1978). The sedimentologic description considered the rock texture, geometry, primary sedimentary structures, and fossil content (*e.g.*, Reading, 1996).

RESULTS

Sedimentary facies

The sedimentary succession exposed in the study area comprises sandy to conglomerate beds (Figure 3), characterized as clast supported conglomerate (Cg), gravel to coarse-grained sandstone (Sg), and very fine- to fine grained sandstone (Sh) facies (see more in Table 1; Figures 4A–B). Most of the succession is represented by paleosols (facies P)

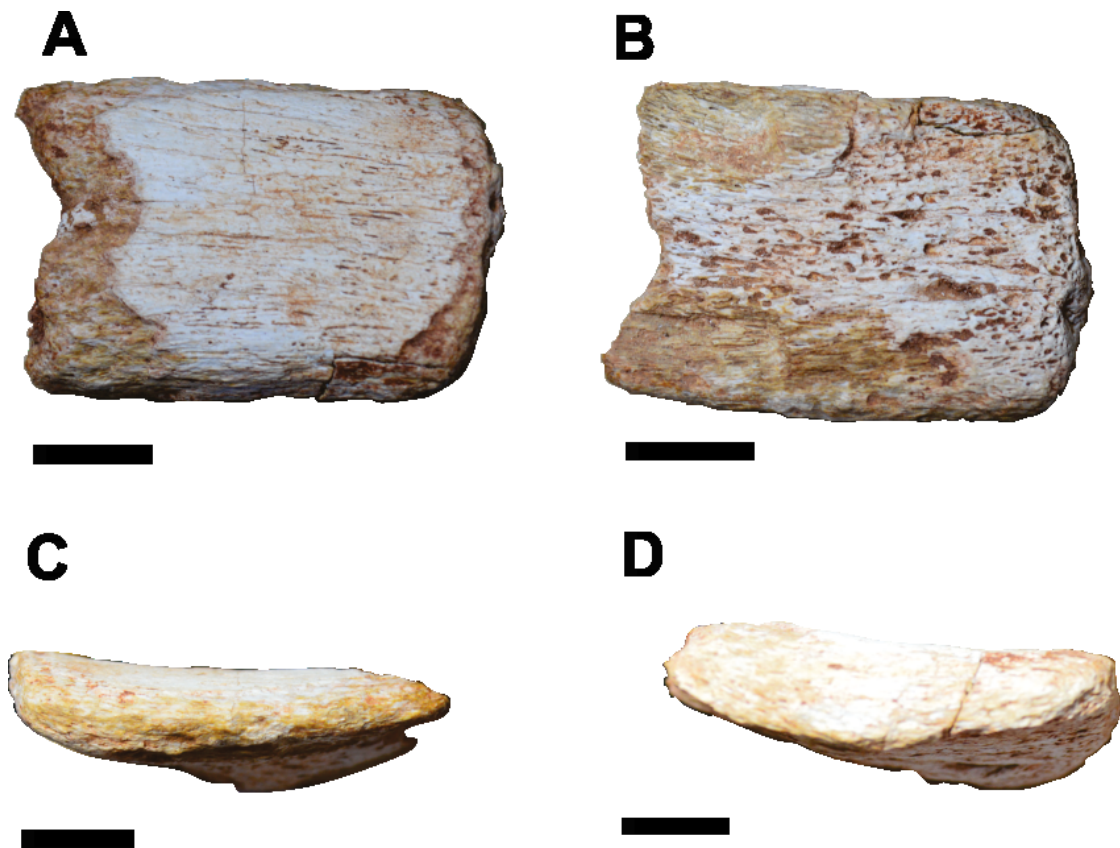


Figure 2. V-0126, a bone remain collected inside the paleoburrow. **A**, smooth surface view. **B**, trabecular surface view. **C–D**, weathered bone edge in lateral views.

Table 1. Sedimentary facies and architectural elements description.

Facies code	Description	Trace fossils	Architectural element
Cg	Clast supported conglomerate, polymictic, with pebbles up to 6 cm in fining upward trend	—	Channels (CH)
Sg	Gravel to coarse-grained sandstone with isolated pebbles in the matrix up to 5 cm in fining upward trend	—	Channels (CH)
Sh	Very fine- to fine grained sandstone with incipient horizontal stratification, locally with low angle cross-stratification	—	Amalgamated sandbars (SB)
Sm	Fine- to coarse-grained sandstone, massive, with carbonated cement, locally with carbonate nodules and irregular basal contact	Rhizoliths	Floodplain fines (OF)
P	Paleosoil with sandy composition and carbonated matrix, with block or massive structures	Rhizoliths, paleoburrow	Floodplain fines (OF)
Fm	Massive siltstone to poorly laminated mudstones	—	Floodplain fines (OF)

associated with sandstone and mudstone (Sm and Fm facies; Figures 4C–D; Table 1). The paleoburrow is preserved in a sandy matrix with pedogenetic signatures (facies P; Figure 4F), associated with rhizoliths mostly represented by root casts and rhizocretions (Figure 4E).

The presence of a channel filled by coarse-grained sediments (Cg, Sg, and Sh facies) interspersing facies with pedogenetic features (P, fine- to coarse-grained sandstone massive- Sm, and massive siltstone- Fm facies) suggests a

“braided-like” fluvial system consisting of low sinuosity, amalgamated single-story channels (*e.g.*, Miall, 2016). The bedload-dominated transport eroded the floodplain deposits and resulted in conglomerates with fining-upward trends (McKie, 2011). The amalgamated sandbars represent migration along the floodplains generated by unconfined flows (Fielding *et al.*, 2011). The floodplain fines are associated with sandy-dominated floodplain deposits and subordinated fine-grained sediments. The presence of block structures,

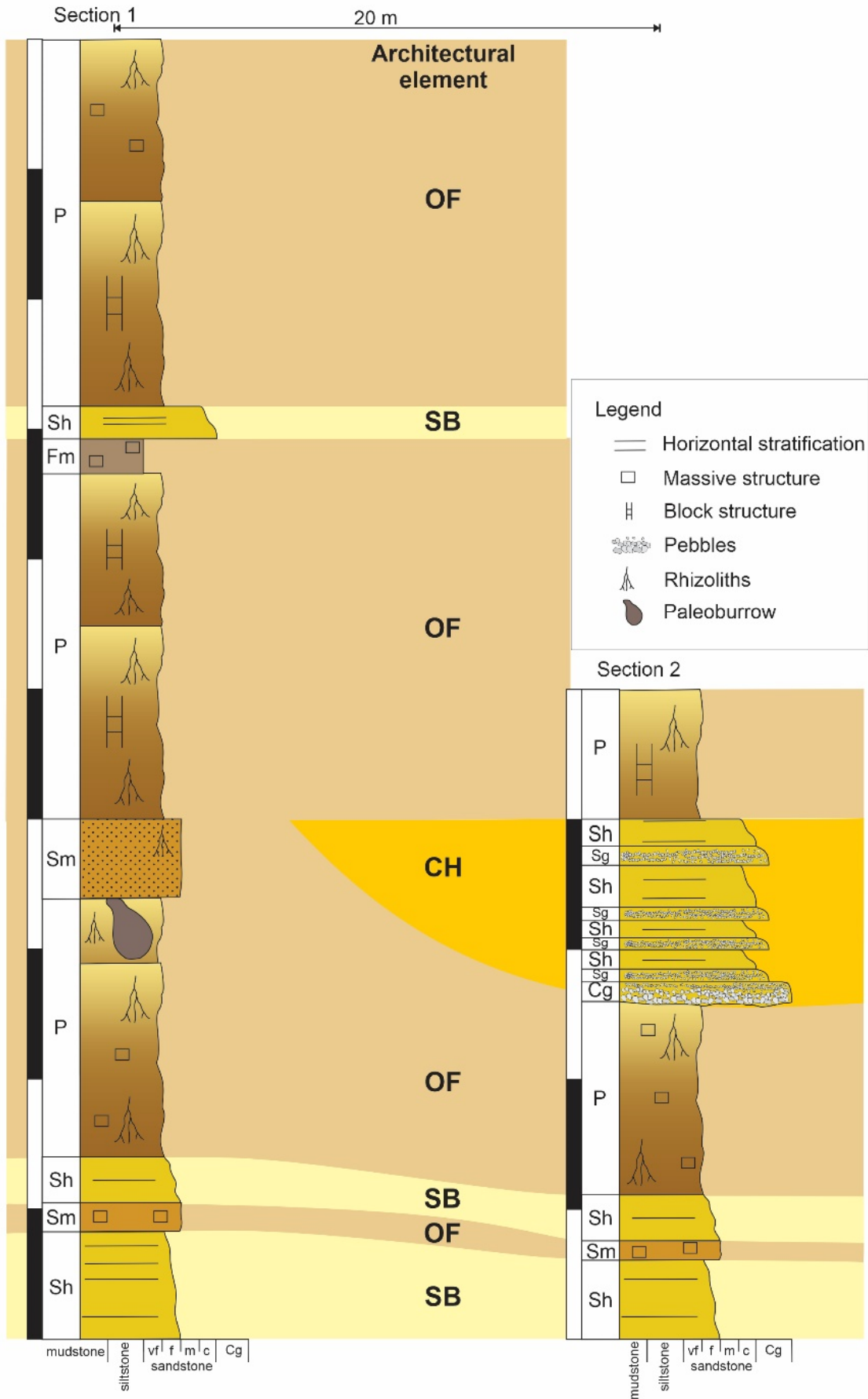


Figure 3. Measured columnar sections with the stratigraphic position of the paleoburrow. Scale bars are in meters.

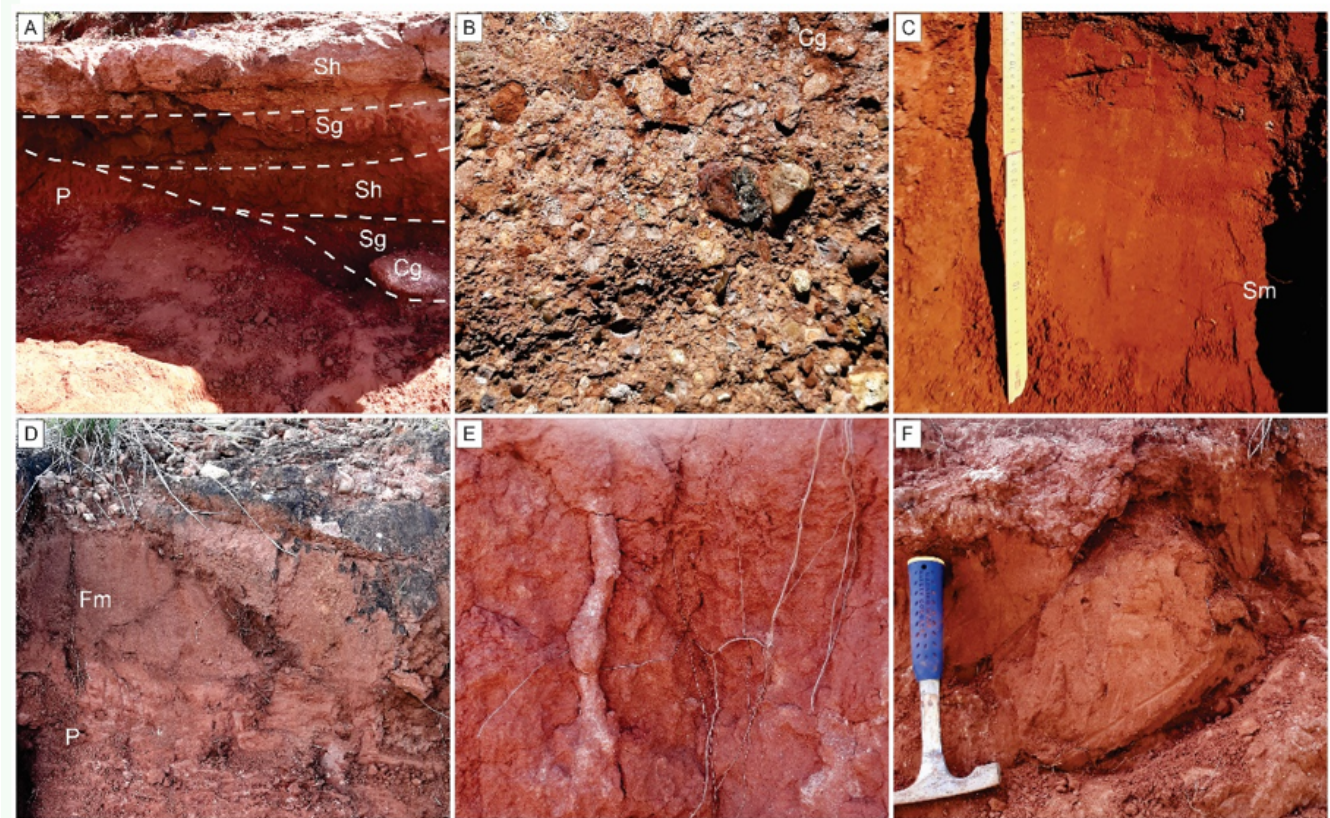


Figure 4. A, the facies are configured as Cg, Sg, and Sh facies, related to a channel filled by coarse-grained sediments and pedogenetic features facies as P, Sm, and Fm. B, clast-supported polymictic conglomerate. C, fine- to coarse-grained massive sandstone with carbonate cement, locally with carbonate nodules and irregular basal contact with rizoliths. D, massive siltstone to poorly laminated mudstone. E, rhizoliths represented by filled roots. F, the studied paleoburrow filled in by a sandy matrix with pedogenetic signatures.

rhizoliths, and the described paleoburrow are evidence of subaerial exposition (Hajek & Edmonds, 2014). The identified architectural elements (channels, amalgamated sandbars, and floodplain fines) are typical for Adamantina Formation (*e.g.*, Batezelli, 2017).

Burrow description

The burrow was exposed in a longitudinal section at the outcrop (Figure 4F). It presents a simple, flask morphology, with a narrow upper shaft connecting a lower, wider chamber. The paleoburrow is 35.39 cm deep and 32 cm wide in its widest axis, thus the depth/width ratio is 1.1. It has a subvertical orientation and a simple architecture, lacking adjacent chambers or branches. The burrow neck is almost perpendicular to the sediment layers, while the chamber is directed $\sim 40^\circ$, with its final portion not very wide (Figures 5A–B). The contact between the burrow and the host rock is sharp, but lacks ornamentation (*e.g.*, claw- or tooth-made scratches). The burrow is passively filled by the sediments that compose the overlying bed (Sm facies). Additionally, scattered fragments of tetrapod bones were found within the filling, more precisely above the level in which the burrow reaches its wider diameter. The larger bioclast fragment is 4 cm long, 2.7 cm wide, and

0.9 cm thick (code LAPE-V-0126). The cracks parallel to the main axis in LAPE-V-0126 indicate the Weathering Stage 1 of Behrensmeier (1978), while smaller fragments probably fall into advanced stages. No secondary bioturbation was found within the burrow filling.

DISCUSSION

On the origin of the structure

As presented in the results, the studied material is an incomplete burrow, showing a simple architecture, with a flask shape and a narrow upper shaft connecting a lower, wider expansion. The burrow lacks diagnostic features (such as the presence of surface ornamentation or bilobate floor (Cardonatto & Melchor, 2021) that allows recognizing a particular producer. Thus, phenetic and coincident correlations (following Carrano & Wilson, 2001) are necessary to tentatively infer a tracemaker among the vertebrate groups present in the Adamantina Formation (mainly crocodylians and dinosaurs) and the Late Cretaceous units of the Bauru Group, such as lungfishes and turtles (see Alves *et al.*, 2013, Martinelli & Teixeira, 2015; Silva *et al.*, 2022). Anurans, lepidosaurs, and mammals were excluded *a priori* from

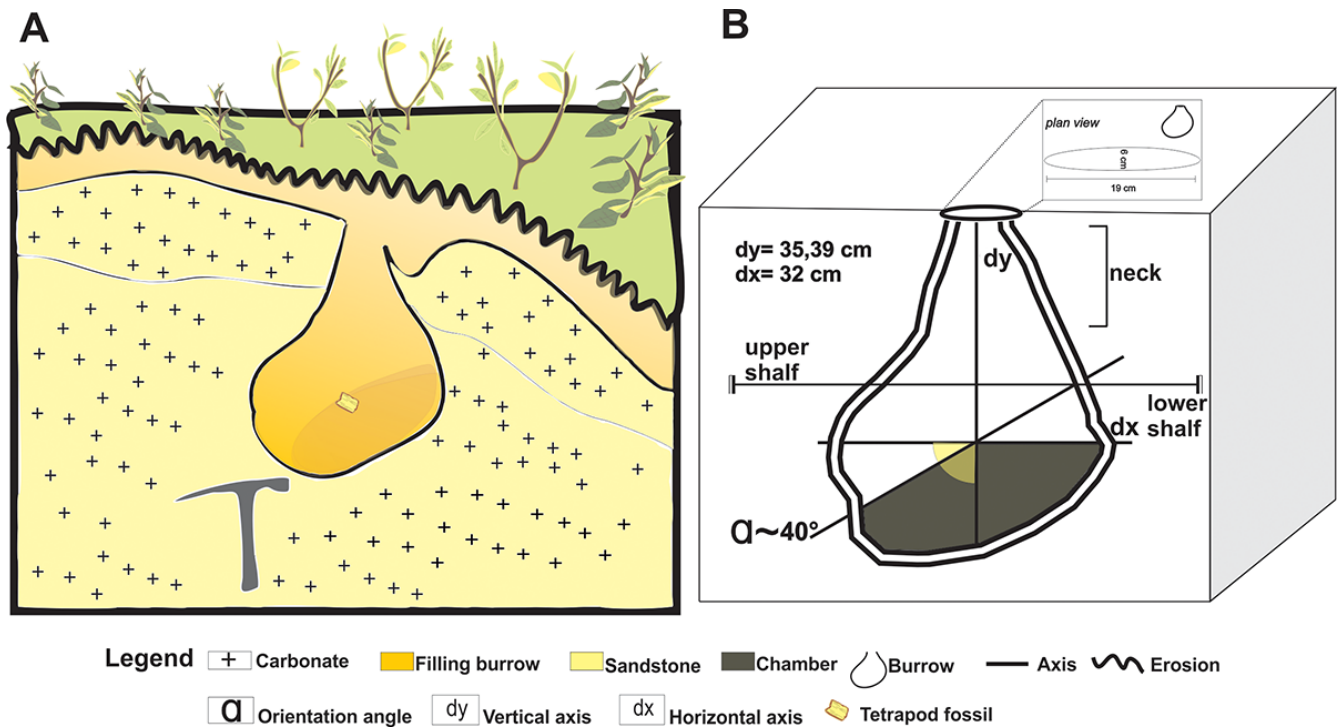


Figure 5. A, sketch of the paleoburrow found and demonstrating the delimitation of the structure, as well as the layers with carbonate nodules, with hammer as 13 cm scale. B, the measurements of the paleoburrow and its orientation angle.

the comparison due the lack of sufficient data regarding potential fossorial habits on the Bauru Group specimens, the controversial taxonomy of some materials, and on the fact that the species known from the Bauru Group are too small to fit the producer of the burrow described here.

The tetrapod burrow record is very scarce in Jurassic and Cretaceous deposits. Therefore, we compiled some information regarding the taxa present on the Marília and Adamantina formations, as well as in other Cretaceous units, to discuss the potential tracemakers. A diverse vertebrate fauna dominated by tetrapods has been described for the Adamantina Formation (see Martinelli *et al.*, 2019). Among them, Mesoeucrocodylia is the most diverse group, comprising sphagesaurid, baurusuchid, and other notosuchian species. Martinelli *et al.* (2019) described a filled terminal burrow in the Campanian–Maastrichtian deposits of the Serra da Galga Formation (formerly Serra da Galga Member of the Marília Formation), assigning it to a tetrapod, possibly to an advanced notosuchian. This burrow is circa 86 cm longer than the one described herein and presents an expanded terminal chamber with an oblique floor ($\sim 30^\circ$), which is not observed in the described paleoburrow.

Even though these structures are likely different in architecture, both are relatively incomplete, hampering deeper comparisons. Other fossil crocodylomorph burrows are unknown so far, precluding a proper comparison. However, some potential fossorial species occur in the crocodylomorphs preserved in the Bauru Group (e.g., Marinho & Carvalho, 2009; Martinelli *et al.*, 2019). Thus, they cannot be ruled out as potential producers of the described paleoburrow.

A simple and sinuous tunnel with a terminal chamber preserved in Cenomanian deposits in the USA was attributed to the euornithopod *Oryctodromeus cubicularis* due to skeletal remains preserved within the burrow filling (Varricchio *et al.*, 2007; Fearon & Varricchio, 2016). Similar material was described from the Albian Otway Group of Australia (Martin, 2009). Despite the burrow architecture and the presence of skeletal remains in the burrow filling being comparable with the paleoburrow found in the Adamantina Formation, the morphology, orientation, and diameter of both burrows are not similar. Moreover, ornithischians were not found in the Adamantina Formation yet, and none of the dinosaur taxa known in this unit (sauropods and theropods) present osteological anatomy that suggest adaptation for deep excavation or fossorial life (Martinelli & Teixeira, 2015; Martinelli *et al.*, 2019).

Silva *et al.* (2022) described several burrows from the Adamantina Formation in São Paulo State, assigning them to terrapins, mainly Podocnemididae. They differ from the burrow described herein by being simple tunnels, with flat floors, low inclination (22°) and lack of chambers or expansions (Silva *et al.*, 2022).

Modern turtles can dig a wide range of burrow morphologies. The most striking example is the Gopher tortoise (*Gopherus polyphemus*) burrows, which comprise simple, unbranched tunnels, with a low-angle ramp (Kinlaw & Grasmueck, 2012), a distinct architecture from that described herein. Regarding modern turtle nests, they are actively filled by the female after the oviposition (also differing from our material) (Landers *et al.*, 1980), and no exception for that is

known from the fossil record so far. Based on these data, we also exclude turtles as the potential producers of the burrow found at Serra do Corpo Seco.

Lungfishes are known for producing vertical aestivation burrows since the Devonian (Hasiotis *et al.*, 1993; Friedman & Daeschler, 2006; Gobetz & Andallanjerner, 2006; Jones & Hasiotis, 2018). Although Paleozoic lungfish burrows show simple, nearly vertical, unbranched architectures and relatively smooth surficial morphologies, the Mesozoic and Cenozoic record often present terminal and wider chambers, where the individual remains coiled inside a mucous cocoon during aestivation period (Hasiotis *et al.*, 1993, 2007). The aestivation burrows of modern lungfishes generally present vertical shafts and slight angle (Greenwood, 1986). Thus, due to the simple morphology, absence of ornamentation, passive filling, and preservation in a flooding plain close to the main channel, we tentatively attributed our burrow to a lungfish excavation.

In summary, the burrow described here present an ambiguous morphology, making it difficult to assign it to a precise tracemaker. Therefore, for the moment, we assign it to a lungfish, and exclude most of the vertebrate taxa known from the Adamantina Formation because of (i) their size (sauropod and theropod dinosaurs and some crocodylian taxa); and (ii) the lack of evidence based on corporal or ichnological/taphonomic data of the group (crocodylians and turtles). New discoveries on the region are necessary to attest this attribution. Besides the inaccurate assignment, it is important to notice that the material described here enlarges the knowledge on the scarce Cretaceous record of paleoburrows in continental settings.

Paleoenvironmental inferences

A vertebrate filled burrow can not only be associated to biological aspects of its producer, but also to the paleoenvironmental conditions in which it was produced and filled, being a very useful tool for paleoenvironmental and paleoecological interpretations. Some burrows can be produced by primary diggers, that are organisms that modify the excavation in secondary way, and the organism just dwelling the burrow (Kinlaw, 1999). These excavations can be made by claws, teeth, and horns of organism in action on substrate (Hopkins, 2005; Hasiotis *et al.*, 2007; Becerra *et al.*, 2013; Hildebrand, 2013). This burrow does not preserve excavation marks caused by claws, teeth, or horns on the inner wall, possibly due to excessive exposure of the structure to environmental conditions, a fact suggested by concretions in the surrounding host rock, showing an exposure time before and during the sediment deposition. The erosive contact influences the overall preservation of the paleoburrow, due to the impossibility of visualizing the continuity of the structure towards the surface or even into the host rock.

The open burrow worked as a trap for the deposition of sediments transported by energy flows of different intensities. The trap is configured by various ways of filling by sediments that vary according to flow velocities, sediment supply, dry or water-filled conditions, and initial placement and

weight of fossil/bones or grains, resulting in homogeneous or stratified filling (Woodruff & Varricchio, 2011; Gallois, 2013; McLoughlin *et al.*, 2020). The fill flow previously carried the fossil fragments into the burrow. The presence of carbonate fills in some cracks and carbonate cement inside some burrows is used to establish further stratigraphic controls on their genesis (Jäger & Heinrich, 1975). These are closely associated with paleosol formation and, therefore, alter the microstructure and ancient edaphic zones (Sycheva *et al.*, 2019). The krotovinas can be filled with ferromagnetic minerals, which are used to paleomagnetic measurements, such as clayey paleosol (Rico & Bidegain, 2013). These structures also indicate a significant biological activity, since it can be filled not only by sediments, but also by coprolite and the remains of other organisms (Pietsch, 2013; Popova, 2015; Barczy *et al.*, 2020). Some digs serve as a natural source of habitat and heterogeneity for the ecosystem, altering the topography, the availability of nutrients for vegetation or even the habitat for aquatic organisms, such as what happens with the extant alligator holes in the region of the Everglades (Palmer & Mazzotti, 2004). Therefore, the capability of a krotovina or burrow to record temporal information or even condition small living environments for animals and plants is relevant, as evidenced in this outcrop, where nodules are found scattered throughout the area that borders the burrow or even the rhizotubules that meet its external wall, not entering the filled part, anchoring a plant on substrate.

Eventually, bone remains found inside a burrow can be associated to its probable producer. However, in our case, the bone remains found within the burrow filling are highly fragmented and somewhat weathered (Figure 6D), the larger bone presenting cracking lines parallel to its main axis (falling into the Weathering Stage 1 of Behrensmeyer, 1978), while the smaller fragments are millimetric and highly decayed. The fragments were found in the middle part of the chamber (Figures 6B–C) surrounded by the homogenous structureless reddish sandstone, indicating a passive burrow filling. The evidence for abrasion and weathering suggests that they had significant residence time outside the burrow previously to the transportation to within the burrow (at least 0–3 years, considering the Weathering Stage 1 of Behrensmeyer, 1978).

In general, bone remains found inside a burrow can be considered as belonging to its producer or occupant only if they are nearly complete, articulated or disarticulated, but with elements still closely associated (*e.g.*, Smith, 1987; Groenewald *et al.*, 2001; Damiani *et al.*, 2003). Therefore, it is unlike that the bone fragments found inside the burrow belong to the producer and denotes an allochthonous origin. Thus, the burrow and the bones had a time of subaerial exposure, a fact suggested by the Behrensmeyer stage, by the erosive contact, by the concretions in host rock and by the non-preservation of excavation marks on the inner wall of the paleoburrow. Paleoburrows help in the interpretation of paleosols, and geologic studies should focus on the linkage between trace fossils and paleosols, and how they can be used to interpret the paleoenvironmental parameters under which those paleosols were formed (Hasiotis & Halfen, 2010).



Figure 6. A-B, burrow chamber filled with massive sandstone. C, chamber delimitation. D, bone fragment found in the middle of the chamber.

Some burrows are passively filled, in the manner of sediment traps, simply by gravity-induced deposition; the fillings may be virtually identical to the host matrix or may be more akin to an overlying layer of sediment (Frey & Pemberton, 1985). The filling of the chamber is a continuation of the top layer; therefore, it exposes lithological composition and similar sedimentary structures, with no evidence of apparent carbonate concretions (Figures 6A–B).

In pedogenic calcretes, carbonate precipitation occurs mainly in the vadose zone above the water table, specifically in cavities in the host rock, similar to what occurs in underground calcretes (Alonso-Zarza *et al.*, 2012). The water table level can be evidenced by burrows, which indicate a markedly ephemeral deposition at the central point of an active sedimentary basin, an important tool for the reconstruction of hydrological and climatic conditions (Voigt *et al.*, 2011). Therefore, a pedogenic origin is suggested for these calcretes dispersed in the host rock, a phenomenon related to an arid or semi-arid environment (Figure 7). The formation of pedogenic calcretes maintains a close relationship with the vegetation of the area, indicated by the distribution of rhizohaloes and

rhizocreations, and which typify the paleosol (Catena *et al.*, 2017; Nascimento *et al.*, 2021). Deep and abundant rhizohaloes and red matrix indicate good drainage and low water table; yellow rhizohaloes suggest relatively high soil moisture and organic matter content (Kraus & Hasiotis, 2006; Smith *et al.*, 2008). The yellow rhizohaloes are seen on the host rock, visible 15–20 cm from the surface. Limestone rhizotubules are also seen, which are usually found in soils that undergo periodic drying (Kraus & Hasiotis, 2006).

CONCLUDING REMARKS

The paleoburrow reported herein is referred to a lungfish as a potential producer according to the evidence presented. This interpretation is supported by its simple architecture and sub-vertical orientation with expansion of a chamber at the end of the structure. Those results also evidence long periods of dry conditions, as previously suggested for the much-debated Adamantina Formation, interrupted by flooding events as suggested by facies analysis and the presence of a lungfish burrow.

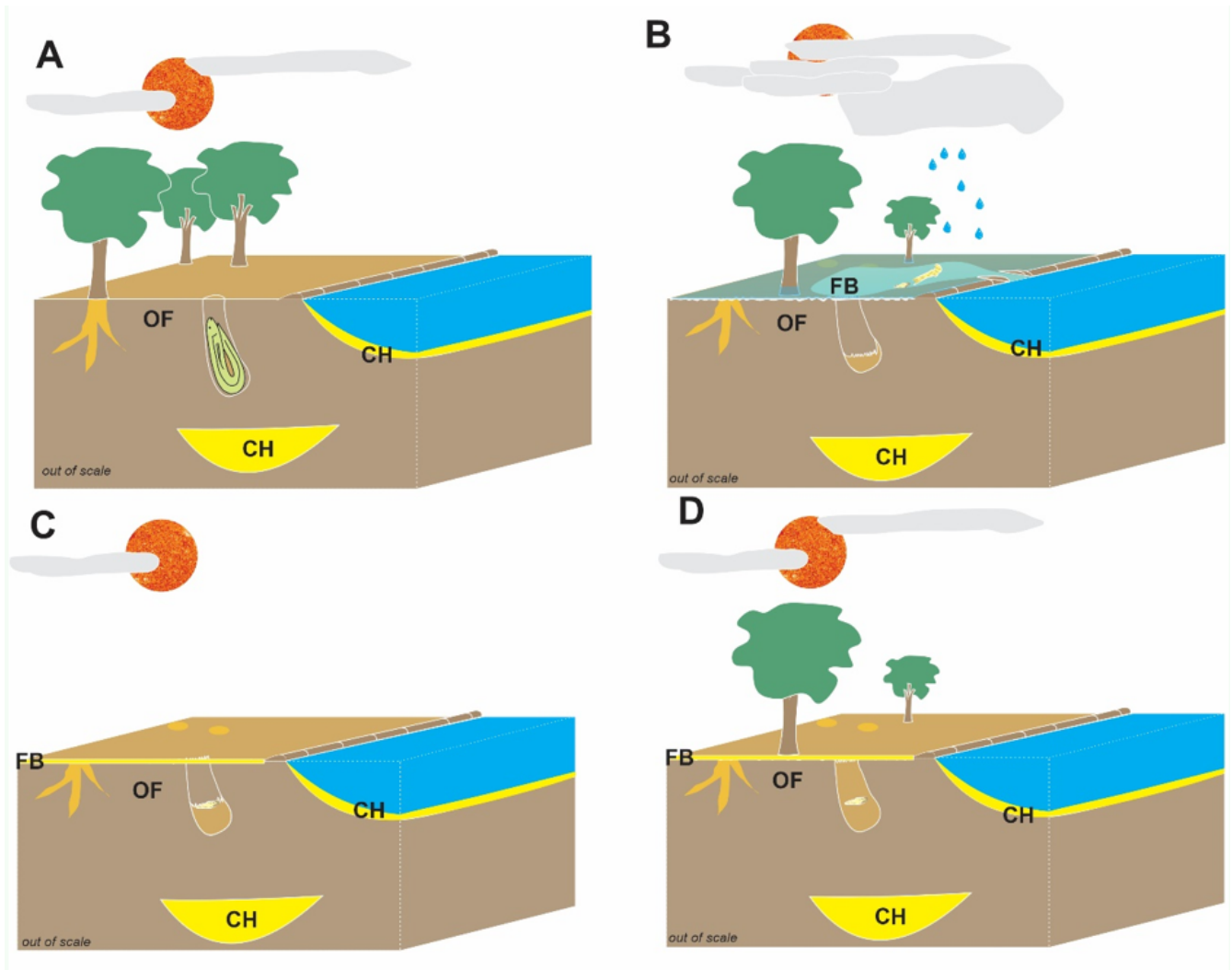


Figure 7. Palaeoenvironmental reconstruction based on paleoburrow, rizhotubules, and depositional model of Adamantina Formation deposits in Serra do Corpo Seco. **A**, construction of a burrow in a flooding plain (overbankfines; **OF**) close to the main channel (**CH**). **B**, flooding events and deposition of amalgamated sandbars (**SB**). **C**, dry periods and slow filling of the opened burrow. **D**, successive flooding events generating filling of the burrow.

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